

1 **Title Page**

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3 **The 1908 Messina Straits earthquake: Cornerstones and the need to step forward**

4

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19 The M 7.1 earthquake that on the 28th Decembre 1908 struck the cities of Messina and Reggio
20 Calabria, facing each other across the Messina Straits, was one of the most destructive events
21 of last century, and one of the deadliest earthquakes in human history, causing a huge death
22 toll, estimated between 60,000 and over 100,000 (Bertolaso et al., 2008). None of the large
23 earthquakes that occurred in Europe in the XX century, some of them with larger magnitude
24 (Grünthal and Wahlström, 2012), resulted to be as deadly as the Messina earthquake, which
25 sparked a wave of solidarity and grief throughout Italy and also worldwide, with the launch of
26 international relief efforts.

27 The earthquake was soon followed by large tsunami waves that swept the coast of eastern
28 Sicily, reaching as far south as the island of Malta, and caused about 2000 casualties (e.g.,
29 Schambach et al. 2020). The very early observations of Fusakichi Omori, that visited the
30 devastated lands two months after the earthquake, suggested an offshore source for both the
31 earthquake and tsunami (Omori, 1909; Valensise, 2019).

32 Over the years many scientists have tried to identify the fault that originated the 1908
33 Messina Straits earthquake. The analyses accomplished so far identified some important
34 constraints for the causative fault, but no comprehensive source model has been presented
35 yet and no general consensus has been reached.

36 Here, we briefly review some of the proposed interpretations of the 1908 Messina fault,
37 commenting on their reliability, and suggest that looking for a different approach would be
38 advisable.

39 Besides the scientific interest and the importance for hazard assessment, the understanding
40 of the origin of the 1908 Messina Straits earthquake has also relevance for the project of a
41 bridge crossing the strait, which always hovers. The bridge projects has its roots back in time,
42 even before the 1908 earthquake, but since the 1970s this issue has been cyclically in political

43 propaganda during the (far too many) election campaigns, in those occasions being
44 considered a necessary infrastructure.

45

46 **Seismology and Geodesy**

47 The analysis of the seismological and geodetic data suffered from the limited (by modern
48 standards) instrumental networks of the beginning of 20th century, and the subsequent
49 record of seismicity has been characterized (luckily) only by moderate-small, sparse,
50 earthquakes that brought little information on major fault planes. Nevertheless, the attempts
51 to analyze the 1908 seismological and geodetic data have produced a considerable number of
52 papers, from the late '80s to the beginning of 2000 (see Pino et al., 2009 for a review), which
53 presented robust results at least for some specific characteristics of the rupture for the 1908
54 earthquake. Also, recent analyses of the contemporary seismicity and of the surface
55 deformation data provide valuable information on the seismogenic layer, the crustal structure,
56 and the stress field in the Straits area (Neri et al., 2021).

57 *Source characteristics of the 1908 earthquake*

58 The investigations of the historical data proposed different fault planes, derived from the
59 inversion of either the first motion polarities or the levelling lines, or from both. Most of these
60 models share an extensional focal mechanism, on a plane roughly oriented in the N-S
61 direction (N345°-N15°), with similar low angle dip (30-40°) to the east (Fig. 1). The well-
62 assessed evidence of northward rupture propagation, constrained from the analysis of the
63 original waveforms (Pino et al., 2000) is fully consistent with the above indications and also
64 confirms the location of the maximum strain energy release in the Straits area. It is worth
65 noting that the location and the orientation of the fault plane, and the rupture propagation
66 direction also appear to be consistent with the simulation of the felt intensity reports, through

67 full wavefield synthetic seismograms (Convertito and Pino, 2014), which also indicates
68 preference for a slightly rotated clockwise fault plane.

69 More recently, some authors presented alternative source geometries, characterized by
70 significantly different location, geometry, or kinematics (e.g., Aloisi et al., 2013; Meschis et al.,
71 2019; Barreca et al., 2021). However, the results of these studies rely on *ad hoc* hypotheses,
72 derived from the levelling data alone, either neglecting the seismological data (e.g., first
73 motion polarities) or not accounting for the well-assessed constraints derived from the
74 seismological analyses, and some of them also have inconsistencies in the analysis (De Natale
75 and Pino, 2014; Argnani, 2021; Pino et al., 2021). It is helpful to recall here that the
76 geographical distribution of the first motion polarities is not wide enough to univocally
77 determine the focal mechanism, however the net separation of compressions and dilations in
78 the NE quadrant constrains the W-dipping plane to have azimuth roughly in the range $N0^\circ$ -
79 30° and dip around 50° (e.g., Capuano et al., 1988).

80 The re-examination of the first arrival times provided remarkably stable results for the
81 epicentral location, indicating that the rupture started in the SE area of the Straits (Michelini
82 et al., 2005), in agreement with both the location of the fault plane in the Straits area and its
83 approximately NS orientation, and with the northward propagation of the rupture as well.

84 As for the fault extension, it seems to be confined along a 40-45 km length, as results for
85 several models (Fig. 1). Some authors also derive dislocation outside the Straits area (e.g.,
86 Amoruso et al., 2002). In principle this cannot be excluded, but this result is inconsistent with
87 the slip linear distribution derived from the seismograms' analysis (Pino et al., 2009).

88 *Present seismicity and strain in the Straits area*

89 During the last decades, in the Straits only minor earthquakes occurred, defining a cloud of
90 scarce and sparse seismicity, characterized by $M \leq 4.0$ and depth in the range 5-15 km, with

91 nearly all the $M \geq 3.0$ events occurring between 10 km and 15 km (e.g., Scarfi et al., 2009; Neri
92 et al., 2021).

93 Tomographic images derived from the analysis of these events are not very well defined,
94 generally evidencing the presence of a band of negative anomaly of the P-wave velocity,
95 running along the Straits axis (e.g., Scarfi et al., 2009), while the available focal mechanisms –
96 derived either from polarities (Scarfi et al., 2009) of the first arrivals or from waveform
97 inversion (Neri et al., 2021) – are associated with ESE-WNW seismic strain direction. This
98 latter evidence is fully consistent with the present crustal strain determined from GPS
99 analysis, which display extension normal to the coast Sicilian side of the Straits with
100 maximum rate of ~ 65 nanostrains/yr (Serpelloni et al., 2010).

101

102 **Geology**

103 In the geological approach it has been inferred that the 1908 Messina Straits earthquake, due
104 to its great size, left a trace in the geological record, breaking a fault plan that can be imaged in
105 the subsurface and which has possibly ruptured the seafloor. The Messina Strait is a
106 relatively small area, that can fit within a 20 x 40 km box. The bathymetric data and
107 multichannel reflection seismic profiles acquired in the last 20 years have reached a
108 reasonable coverage, though deep seismic data are still lacking, mostly because of the
109 difficulties of towing long streamers in such a narrow stretch of sea. These data allow a
110 reasonable image of the seafloor morphology (e.g., Ridente et al., 2014; EMODnet, 2016) and
111 of the sediments and rocks for a few km below it (e.g., Argnani et al., 2009; Argnani, 2011;
112 Argnani 2021; Argnani, 2022; Barreca et al 2021). Nevertheless, these data have been often
113 undervalued or forced to support preconceived interpretations.

114 Recent attempts to identify the surface rupture of the fault responsible for the 1908 Messina-
115 Straits earthquake (Meschis et al., 2019; Barreca et al., 2021) have proved poorly reliable
116 (Pino et al., 2021; Argnani, 2022; Pino, 2022), in spite of echoes in the press. The high angle
117 Messina-Taormina Fault (MTF; Fig. 1) has been proposed by Meschis et al (2019) on the basis
118 of inversion of geodetic data as rupturing the seafloor for a length of ca. 60 km. This fault
119 however is not visible in any of the seismic profiles crossing the Sicilian slope (Argnani et al.,
120 2009; Barreca et al., 2021) and is not consistent with the first motion polarities observed at
121 seismic stations, whatever crustal hypocentral depth is assumed (Capuano et al., 1988). The
122 W-Fault, mostly running along the Messina channel, has been traced by Barreca et al (2021)
123 using high resolution seismic profiles. The evidence for this fault, however, is not particularly
124 convincing, with the data being over-interpreted (Argnani, 2022). The recent activity of this
125 fault is doubtful, as its fault scarps are mostly erosional remains at seafloor, and elsewhere the
126 fault is sealed by Late Pleistocene sediments (Argnani et al., 2009; Argnani 2011, 2022). This
127 fault appears related to the fault system present onshore in Calabria, the age of which span
128 from Late Pliocene to Middle Pleistocene (Selli, 1978; Ghisetti, 1992). These age constraints
129 suggest that the 1908 Messina fault is active within a tectonic regime that has operated in the
130 last 120 kyr. Besides that, the W-Fault has some critical seismological and geodetic issues, like
131 the downward cutoff of seismicity, or the assumption that a long-term creep along the fault
132 preceded the earthquake (Neri et al., 2021; Pino et al., 2021; Pino, 2022).

133 Given the constant interest for the 1908 earthquake and the uncertainties associated to the
134 issue, other faults located in the surrounding of the Messina Strait have been proposed in the
135 literature as responsible for this event, often on the basis of regional tectonic considerations,
136 but without supporting documentation. One example is the Ionian Fault, which was originally
137 traced along the western flank of the southern portion of the Messina channel (Fig. 1; Polonia
138 et al., 2012), suggesting that "immediately south of the Messina Straits, a steep east-dipping

139 fault is present at depth and satisfies the requirements of the causative fault for the 1908
140 earthquake." Later, on the basis of regional considerations, the same group of authors (SgROI
141 et al., 2021) changed their view and proposed a tectonic arrangement where right-lateral
142 displacement along the Ionian Fault (now with a different trace; Fig. 1) causes extension in
143 the Messina Strait as a kind of deformation at the northwestern tip of the fault (e.g., Kim and
144 Sanderson, 2006). None of the two hypotheses satisfy the seismological and geological
145 evidence. In the first case no fault has been imaged in the southern stretch of the Messina
146 Straits (Argnani et al., 2009, Barreca et al., 2021) and the proposed fault ignore the northward
147 rupture propagation inferred by seismology (Pino et al., 2000). The second one appears as an
148 ad-hoc model, and the occurrence of a northwestern tip in the Ionian fault contrasts with the
149 hypothesis, supported by the authors, that this fault is lithospheric and propagates from the
150 hinterland into the Ionian foreland.

151 Seismic reflection data show that there is no large N-S fault cutting the Pleistocene
152 sedimentary successions south of the epicenter of Schick (1977), and that in the northern part
153 of the strait the faults imaged by seismic profiles are too short to account for a large
154 earthquake (e.g., Argnani et al., 2009), leaving aside that these faults could be inactive.

155 Although seismic reflection data do not offer an image of the fault plane responsible for the
156 1908 Messina Straits earthquake, they still provide useful constraints on this fault, which has
157 to be blind and not rupturing the uppermost 2-3 km below the seafloor.

158

159 **Outlook and Closing Remarks**

160 The limited seismological and geodetic data for the 1908 Messina Straits earthquake allow to
161 constrain a low-angle, ca. N-S-trending and E-dipping fault, with a maximum length of ca. 40-
162 45 km, and a northward propagation of the rupture from a depth of 8-10 km (e.g., Pino et al.

163 2009). Shallow seismic profiles indicate an indirect dating of the origin of the 1908 Messina
164 fault and allow to reject the occurrence of a fault plane rupturing the seafloor. Therefore,
165 geological and seismological data both point to a blind fault. In this respect a re-examination
166 of the morpho-bathymetry of the Messina Strait is advisable, with a closer inspection at the
167 morphology of seabed looking for signs of the occurrence of an active, blind extensional fault.

168 The acquisition of a deep seismic profile, possibly amphibian, crossing the Messina Strait
169 could be a way to image the plane of the blind fault, although the tectonic complexity of the
170 Calabrian Arc may hinder the result. In the event a fault plane is imaged in the deep
171 subsurface, with an upper tip likely deeper than 3-4 km, it can be verified whether some sort
172 of deformation connected to the fault is observed in the shallow seismic profiles so far
173 acquired. In this case it would be possible to transfer the information to the grid of seismic
174 profiles, thus characterizing the geometry of the fault and better understanding its evolution.
175 Should new and successful geophysical acquisitions be made in the near future, the
176 seismogenic fault plane could also become a possible target for an ICDP drilling aimed at
177 intercepting and monitoring an active fault.

178 In any case, we note that any new interpretation cannot be based on a partial analysis or on
179 arbitrary assumptions, and that the proposed model must take into account all reliable
180 constraints based on available data and observations.

181

182 **Data and Resources**

183 No data or additional sources were used in writing this article

184

185 **Declaration of Competing Interests**

186 The authors acknowledge there are no conflicts of interest recorded.

187

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300 **List of Figure captions**

301 **Figure 1.** Morpho-bathymetry of the Messina Strait region with superimposed the trace of the
302 faults inferred by Meschis et al. (MTF) and Barreca et al (W-Fault). The Ionian Fault-Messina
303 Strait Fault system (MSF) of Polonia et al. (2012) and the Ionian fault (IF) of Sgroi et al. (2021)
304 are also shown; for this latter fault the greenish and purple fields represent areas subject to
305 extension and compression, respectively. The area of strongest shaking produced by the 1908
306 earthquake is outlined by a dotted blue line. The black box is the surface projection of the
307 model fault proposed by Boschi et al. (1989). The small yellow star indicates the epicentral
308 location of Schick (1977). The white star locates the epicenter proposed by Michelini et al.
309 (2005), and the white arrow indicates the rupture directivity (Pino et al., 2000). The thick
310 white bars show the geodetic strain (Serpelloni et al., 2010). ME= Messina and RC= Reggio
311 Calabria Bathymetric data have been obtained from EMODnet (2016) and topography is from
312 Shuttle Radar Topography Mission (Farr et al., 2007).

313

Figure 1 colour

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